

northern side, though this is probably an illusion, occasioned by the difference in relative altitude between Särtäng and the Gobi. There do not appear to be any similar outstanding peaks to the west of the culminating knot of Anambaruin-ula. Such peaks as do exist in that direction bear streaks of snow only, not continuous snowfields, such as might give rise to rudimentary glaciers. From the *firn* expanses of the great central knot certain small hilly »ice-braes» do indeed proceed, yet they too appear to be of a rudimentary character.



Fig. 294. ON THE SIDE-TERRACE OF A GLEN.

The next transverse glen of the first magnitude is Tsagan-tschiloto, though it is not so big as either Aksä or Dschong-duntsa. From it one can count four parallel crests or ridges to the south, although two of these are but branches or offshoots of two larger ranges. Portions of the glens of Ölken-tänesing and Aratänesing are also visible from the same place, though one can only surmise what is the position of the former, namely that it lies behind the dark ridge. Tsagan-tschiloto is formed by at least two glens containing springs, and these unite at the point where we crossed over the system.

After that the slopes of the mountains still continued to be scored by countless numbers of small watercourses and rivulets, tiring to travel over. We nowhere observed hard rock, nothing but gravel-and-shingle detritus, grey granite, striped granite, gneiss, several varieties of mica-schist, quartzite, crystalline schists, greenstone etc.; fragments of stone of rather large dimensions were not at all uncommon. In the larger transverse glens we generally saw, about a couple of kilometers higher up than our route, very steep mountains of a red colour; judging from their shape they consisted for the most part of the accumulated products of disintegration. And it is from this circumstance that the next large transverse glen derives