

narrow valley floor of the present stream. Evidence of similar terracing in the range southwest of Askhabad will be given in a later section. We saw no indications of deltas or other shoreline features here—nothing but the results of forward-washing stream action. It may be that the chief evidence for drawing the Quaternary Caspian shoreline near Kizil-Arvat is to be found in the agreement of the altitude of the plain, a few miles from the mountain base, with the level of the sea as determined by recognizable shorelines elsewhere. In that case its location can, of course, be only approximate.

At Bakharden, between Kizil-Arvat and Askhabad, we rode out to the dunes along the course of a small stream, whose occasional floods keep a graded passage open among the sands for several miles from the mountains. The surface of the

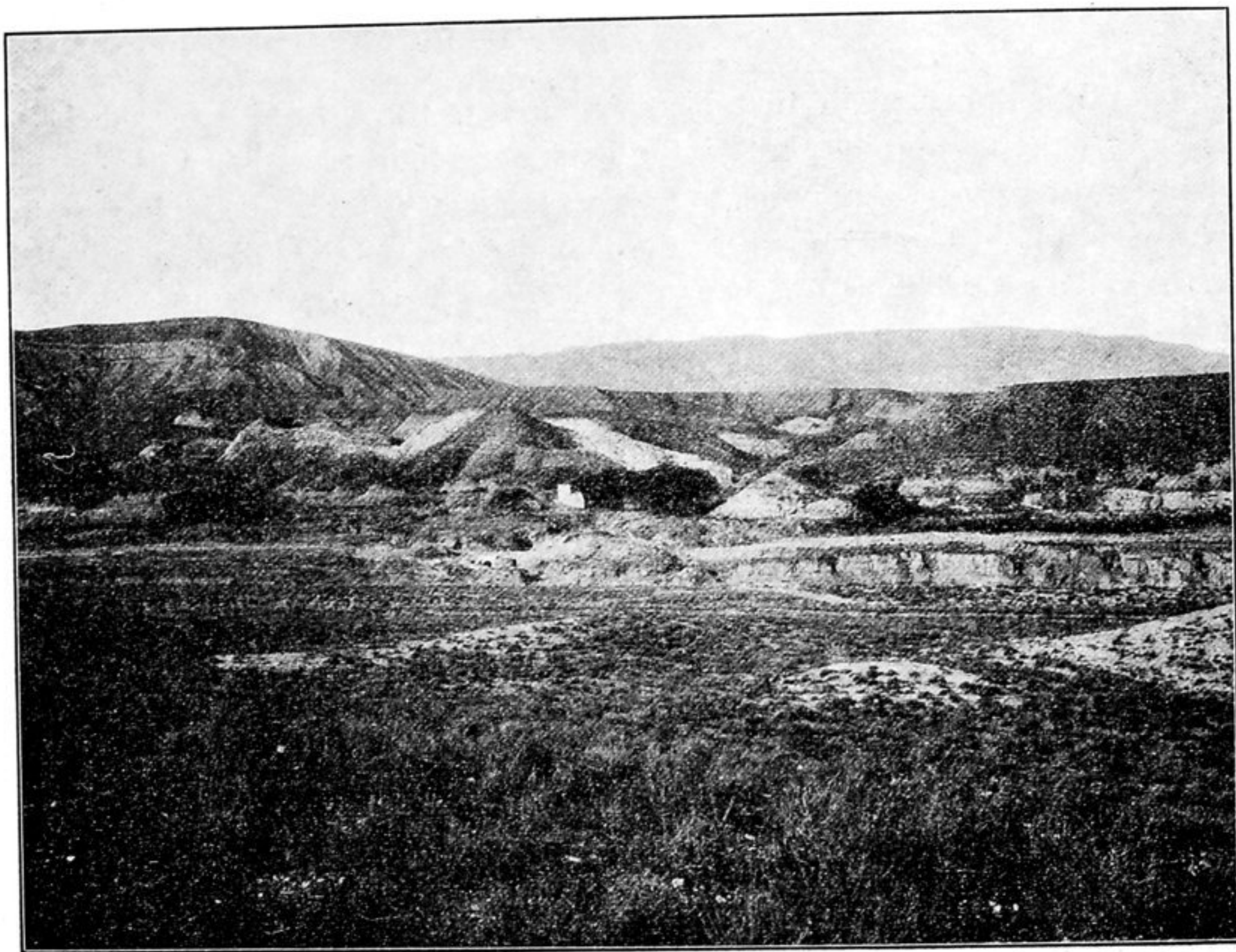


Fig. 22.—Dissected Terraces at the base of the Kopet Dag, south of the Kizil-Arvat, looking southwest. The horizontal limestones of the Mountain on the left are suddenly bent down so as to pass under the clays of the Terraces in the middle distance.

sands was irregular at first (fig. 23); then the dunes began in moderate relief, seldom exceeding 15 feet in height. The scarps of the crescentic dunes or barkhans (fig. 24) were to the west, as if the sands had been drifted by easterly winds. The sand bore a scanty growth of grass, except on the freshest dunes. Between the dunes and the mountains there was no sign of shore-terrace or delta observed. The piedmont slope extends forward without interruption as far as we saw it. It should, of course, be remembered that the abandoned Caspian shorelines, wherever they stand on the piedmont plain, may be faint and not easily recognizable; nevertheless, they were recognized so easily at Baku, Krasnovodsk, and Jebel, that failure to see